

# Ammonite faunas and the 'standard zones' of the Cenomanian to Maastrichtian Stages in their type areas, with some proposals for the definition of the stage boundaries by ammonites.

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The 'standard zones' of the Cenomanian to Maastrichtian Stages as applied to the western European region are critically evaluated, and their development within the type areas of these stages is reviewed. Proposals for defining the bases of the Upper Cretaceous stages worldwide using ammonites are: Cenomanian: appearance of *Hypoturrilites* Dubourdieu, 1953 or *Neostlingoceras* Klinger & Kennedy, 1978; Turonian: appearance of *Pseudaspidoceras flexuosum* Powell, 1963; Coniacian: appearance of *Forresteria* (*Harleites*) *petrocoriensis* (Coquand, 1859); Santonian: appearance of *Texanites* species; Campanian: appearance of either *Placentoceras bidorsatum* (Roemer, 1841), *Submortonoceras* Spath, 1921, or on the basis of evolutionary changes within the *Scaphites hippocrepis* (De Kay, 1827) lineage; Maastrichtian: appearance of *Pachydiscus neubergicus* (Von Hauer, 1858).

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## Introduction

The type areas of the Cenomanian, Turonian, Coniacian, Santonian, Campanian and Maastrichtian Stages provide, to varying degrees, a series of localities where the visitor can see for himself what d'Orbigny, Coquand and Dumont respectively regarded as typical fossils and lithologies of each of these six divisions of the Upper Cretaceous. As an extensive literature demonstrates, they are none of them very suitable for defining the relevant stage boundaries, for with the exception of the Campanian, there are sedimentary breaks around the upper and lower limits of all these stages in their type areas. In spite of this, they remain very suitable reference points for the Cretaceous stratigrapher, because each type area shows a variety of facies and faunas in what are physically and politically readily accessible parts of the world.

In this review I outline the 'standard zones' of the Upper Cretaceous stages that are in current usage, discuss their validity and development in the type areas, and propose what are, in am-

monite terms, likely to be potential marker points for defining the bases of stages.

My own view is that stages are packages of zones, and that it is the base of a zone that is the most sensible way to define the base of a stage.

Discussions at the Copenhagen symposium revolved around appearances of individual taxa on species or genus level, however, and in deference to this my suggestions are made to conform with such an approach.

I have attempted to define boundaries based on ammonites that fall within the breaks at the base of the Upper Cretaceous stages in their type areas (the Campanian excepted, of course). Some of the proposals are unfamiliar, but experience has shown that new, finer biostratigraphic subdivisions, if they are real rather than paper stratigraphy, are often rapidly and widely recognised following their first description. Thus the *Neocardioceras juddii* Zone, introduced by Wright & Kennedy in 1981, was originally documented from southern England, with 3 other records from France and the Germanies. It is now widely recognised in the U.S. Western Interior,

SUBSTAGE	STANDARD ZONES	VARIANTS/SUBZONES
LOWER TURONIAN (part)	<i>Watinoceras coloradoense</i>	
UPPER CENOMANIAN	<i>Neocardioceras juddii</i>	Horizon A
	<i>Metoicoceras geslinianum</i>	Unnamed <i>Nigericeras</i> + <i>Thomasites</i> fauna
	<i>Calycoceras guerangeri</i>	Also called <i>E. pentagonum</i> or <i>C. naviculare</i> Zone; both range above
MIDDLE CENOMANIAN	<i>Acanthoceras jukesbrownei</i>	
	<i>Acanthoceras rhotomagense</i>	<i>Turrilites acutus</i>  <i>Turrilites costatus</i>
LOWER CENOMANIAN	<i>Mantelliceras dixoni</i>	Sometimes called <i>M. orbignyi</i> Zone (a synonym of <i>dixoni</i> )
	<i>Mantelliceras mantelli</i>	<i>Mantelliceras saxbii</i>  <i>Neostlingoceras carcitanense</i>
UPPER ALBIAN (part)	<i>Stoliczkaia dispar</i>	<i>Stoliczkaia dispar</i>

Table 1. Cenomanian ammonite zones.

and can be correlated with the portuguese, Middle East and West African sequences via common elements of both Boreal and Tethyan aspect.

## Cenomanian

When Alcide d'Orbigny began to divide up the Cretaceous system into stages, he at first recognised only two in what is now known as the Upper Cretaceous: Turonien and Sénonien (*Palé-*

*ontologie Française, Terrains Crétacés, II Gastéropodes*, pp. 403–406). With respect to the Turonian, his words are: “je propose de désigner à l'avenir l'étage qui m'occupe sous le nom de terrain TURONIEN, de la Ville de Tours (*Turonnes*) ou de la Touraine (*Turonnia*), situées sur ces terrains” (d'Orbigny 1843, p. 405), defining the Turonian as equivalent to the Craie Chloritée, Craie tuffeau, Glauconie crayeuse, Grès Vert Supérieur etc., and taking the name from Touraine (Roman Turonia). Five years later, he realized

that two distinct ammonite and rudist faunas were present, and he restricted the term Turonian to beds corresponding to his third zone of rudists, yielding "*Ammonites lewesiensis*, *peramplus*, *Vielbancii*, *Woolgari*, *Fleuriausianus*, *Deverianus* etc.", "le plus beau type côtier étant très prononcé dans toute la Touraine, et nous donnerons à la partie inférieure le nom d'étage *Cénomani*, le Mans (Cenomanum), en montrant à la fois le type sous-marin" (d'Orbigny, 1848, p. 270).

In the second volume of the *Prodrome* d'Orbigny (1850) listed 46 ammonites as being characteristic of the Cenomanian, of which 10 were specifically cited from Sarthe, in which Le Mans lies. Localities mentioned are Saint-Calais, La Flèche, Cérans, Ecommoy, Grand Luce, Cou-drecieux, Vibraye, Lamnay and La Ferté-Bernard.

A few ammonites from the Cenomanian of Sarthe were described by Guéranger (1867) in his *Album Paléontologique du département de La Sarthe*, one of the earliest publications with photographs of fossils, and the stratigraphy was investigated in particular by Guillier (1886). The ammonites were then neglected for almost 100 years, until Hancock (1960) listed all the stratigraphically localized Cenomanian ammonites he was able to trace in the Le Mans and Paris Collections, as well as new material collected by him, a total of 161 specimens. This work forms the basis of all subsequent studies.

Latterly, large collections of new material have been assembled, largely through the efforts of Pierre Juignet and André Lebert, so that more than 3000 individuals, accurately placed in sequence, are now available. These have been the basis of preliminary lists in Juignet (1974, 1977), Juignet et al. (1973, 1978) and Kennedy & Juignet (1973); systematic revision is now in progress (Kennedy & Juignet 1983, in press).

In spite of recent opinions that the environs of Le Mans are inadequate and unsuitable as a type area (Marks 1967a, b; Thomel 1973, p. 763), it was an excellent choice. The area is one of diverse facies, and faunas of great richness and variety. There is, however, a regional hiatus at the base of the Cenomanian in most of western Europe (although not recognised to date in Sarthe), while fossils are rare in the lowest beds in the type area.

The zonation of the Cenomanian in western

Europe has progressed erratically because of the lack of adequate taxonomic work. Wright, Kennedy & Hancock (in press) propose a compromise zonation between the work of Hancock (1960), Robaszynski & Amédéo (1980), and Juignet & Kennedy (1976) that can be applied where fine subdivision of the Lower and Middle Cenomanian is not possible. Combined with the work of Wright & Kennedy (1981), this leads to parallel zonations with greater and lesser degrees of precision, as shown in table 1. (Note that the *carcitanensis* and *saxbii*, *costatus* and *acutus* Zones can be treated as subzones of the *mantelli* and *rhotomagense* Zones respectively).

The chief difficulty with Cenomanian zonation has been the search for an index species for what has been, successively, the *naviculare*, *pentagonum* and *guerangeri* Zone as a result of the recognition that the range of the index species extended (in the case of the first two) outside the restricted zone named after them.

In Sarthe, the oldest Cretaceous sediments are Upper Albian (Juignet et al. in press). These are overlain by poorly fossiliferous Lower Cenomanian with *Mantelliceras*, *Schloenbachia*, *Neostlingoceras*, *Hypoturrilites*, *Idiohamites* and other taxa. Above, the diverse and variable facies of the Sarthe yield good assemblages from all but the *juddii* Zone, where only one example of the index species is known.

The heteromorph ammonites of the family Turrititidae provide the best group for defining the base of the Cenomanian, because they have a limited time range but a wide geographic distribution throughout the world and because a continuous phylogenetic series is known from Upper Albian through to Cenomanian: I would propose the appearance of either *Hypoturrilites* Dubourdiou, 1953 or *Neostlingoceras* Klinger and Kennedy, 1978 as defining the base of the stage.

A reference section should be designated in the Monts du Mellègue to Pont du Fahs region of Tunisia, an area that is easily accessible with permanent sections with over 90% exposure. The succession is thick and continuous, ammonites are abundant and include a wide range of taxa, while planktonic microfossils are present and documented. Added to this, the sequence at Djebel Fguira Salah near Pont du Fahs has already been proposed as a Mesogean hypostatotype for the stage (Salaj 1973, 1974; Salaj & Bellier 1978).

The base of the range of both *Hypoturrilites* and *Neostlingoceras* lies within, and probably at the base of the *Hypoturrilites schneegansi* Zone of Dubourdieu (1953, 1956).

## Turonian

The type area of the Turonian Stage is in Touraine, France, between Saumur (on the Loire) and Montrichard (on the Cher) and was so restricted by d'Orbigny in 1852, following the introduction of the stage by the same author in 1843 and its subsequent limitation in 1848 with the introduction, for the lower part of the sequence, of the étage Cénomanién, as already discussed.

There is now a wealth of literature on the Turonian of western Europe and the ammonite faunas and stratigraphy of the region, the stratotype in particular has been minutely discussed (bibliography in Wright & Kennedy 1981, Kennedy, Wright & Hancock 1983, Kennedy et al. 1984 for ammonites; Robaszynski 1982, Alcaydé 1983, Robaszynski et al. 1983 for stratigraphy and other groups).

There is a measure of agreement on the succession of ammonite faunas, but not on the zonal and substage division, as summarised in table 2. There are a number of relevant comments that must be made on this table:

1. The Cenomanian Turonian boundary corresponds to an important oceanic anoxic event (Schlanger et al. in press, with references) and this markedly affected the sedimentary sequence from the *juddii* to basal *nodosoides* Zone.

2. There is widespread condensation and several breaks around the Cenomanian/Turonian boundary in western Europe, so that the sequence is generally thin and the zonal succession locally incomplete.

3. There is an unnamed assemblage between the *geslinianum* and *juddii* Zones as noted previously (table 1).

4. The *Watinoceras coloradoense* Zone is unfortunately named because although *Watinoceras* is common, the nominate subspecies is absent and the base of the zone is drawn at a lower level than the base of the zone of the same name in the U.S.A.

5. The *turonienne*, *kallei* and *ornatissimum*

Zones of French workers are best regarded as locally recognisable units only.

6. *Kamerunoceras turoniense* occurs with *Mammites nodosoides* (Schlüter, 1871) in the Lower Turonian elsewhere in the world.

The chief discrepancy between French and English workers is the positioning of the Middle–Upper Turonian boundary. The argument of Francis Amédéo and his co-workers is that de Grossouvre (1901) and Collignon (1960) regarded *Romaniceras deverianum* as Upper Turonian, and that historical precedent should be followed. This is spurious; de Grossouvre recognised a lower, Saumurien and an upper, Angoumien Substage of the Turonian, the boundary falling in the *woollgari* Zone of the present account (note that de Grossouvre (1901), in table 37 (p. 830), refers the Tuffeau de Saumur to the Angoumien and the Tuffeau de Bourré upwards to the Salmurien, while Collignon at no point refers to Middle Turonian in the paper cited). It is indeed true that Collignon (1939 p. 96) referred *R. deverianum* to the Angoumien, but he again spoke only of Turonien inférieur and supérieur in that work.

All but one of the 'standard' zones of the Anglo-Saxon (and French) schemes can be recognised in the type area of the Turonian; the *coloradoense* Zone by a single specimen, the *nodosoides* Zone by a rather better assemblage and the *woollgari* Zone by abundant material. The *neptuni* Zone is not represented within the stratotype, but the index species occurs rarely in the immediate environs, as at Villedieu-le-Château (Loir-et-Cher), only 45 cm below the base of the Coniacian Craie de Villedieu.

So far as the base of the stage is concerned, Wright & Kennedy (1981) and Kennedy, Wright & Hancock (1983) argue, on admittedly thin historic evidence, that the *juddii* Zone should be taken as Cenomanian, the *coloradoense* Zone as Turonian, and there is a measure of agreement among ammonite workers on this.

Because of the condensed nature of the successions in Europe, the sequences in Texas and New Mexico appear to be the best described to date for defining the boundary in ammonite terms.

The uppermost Cenomanian to basal Turonian zonal sequence recognised by Cobban (this volume) can be correlated with the European standard as shown in table 3.

SUBSTAGE	GALLIC VIEW	ANGLO-SAXON VIEW	SUBSTAGE
LOWER CONIACIAN	<i>Reesideoceras petrocoriense</i>	<i>F. (Harleites) petrocoriensis</i>	LOWER CONIACIAN
UPPER TURONIAN	<i>Subprionocyclus neptuni</i>	<i>Subprionocyclus neptuni</i>	UPPER TURONIAN
MIDDLE TURONIAN	<i>Romaniceras deverianum</i>	<i>Collignoniceras woolgari</i>	MIDDLE TURONIAN
	<i>Romaniceras ornatisimum</i>		
	<i>Romaniceras kallesi</i>		
	<i>Kamerunoceras turoniense</i>		
LOWER TURONIAN	<i>Mammites nodosoides</i>	<i>Mammites nodosoides</i>	LOWER TURONIAN
	<i>Watinoceras coloradoense</i>	<i>Watinoceras coloradoense</i>	
UPPER CENOMANIAN (part)	<i>Neocardioceras juddii</i>	<i>Neocardioceras juddii</i>	UPPER CENOMANIAN
	<i>Metioceras gesinianum</i>	(unnamed <i>Nigericeras-Thomasites</i> fauna) ----- <i>Metioceras gesinianum</i>	(part)

Table 2. Turonian ammonite zones.

U.S.A.	EUROPE
<i>Mammites nodosoides</i>	<i>Mammites nodosoides</i>
<i>Vascoceras birchbyi</i>	<i>Watinoceras coloradoense</i>
<i>Pseudaspidoceras flexuosum</i>	
<i>Neocardioceras juddii</i>	<i>Neocardioceras juddii</i>
<i>Vascoceras n. sp.</i>	(unnamed <i>Thomasites</i> fauna)
<i>Sciponoceras gracile</i>	<i>Metoicoceras geslinianum</i>

Table 3. Uppermost Cenomanian-lowermost Turonian Zones.

*Watinoceras coloradoense* is a *birchbyi* Zone species in the U.S.A. and is represented in Europe by a subspecies different from the nominate one, occurring with other *Watinoceras* and *Thomasites* species.

The base of the *flexuosum* Zone is here proposed as the base of the Turonian. In its type area the zone yields a mixed fauna of ammonite species (Powell 1963) including some that are short-ranging and extend into the Western Interior province, Mexico and California, the Tethyan *Vascoceras*-dominated faunas of east and north Africa and the European boreal realm, while closely allied, if not conspecific forms also occur in South America, Madagascar, Japan and elsewhere.

The type section of the zone is in Calvert Canyon, west Texas, where there is easy access, 90% exposure, abundant ammonites and inoceramids (the latter the *Mytiloides opalensis* of Kauffman, not Böse) while the pelagic facies represented yields planktonic microfossils.

It is arguably unsuitable as a boundary stratotype because the zonal fauna is isolated. Better sections are to be found elsewhere in Texas and New Mexico, where the underlying *juddii* Zone is ammoniferous. I defer to Dr. W. A. Cobban to suggest a boundary section.

This boundary corresponds to the appearance of the *Mytiloides* lineage, which means that the

base of the stage can be easily recognised where ammonites are absent.

If this definition is accepted, much of the 'Lower Turonian' of the Tethyan Realm must be reclassified as Upper Cenomanian, as has been recently confirmed by a number of workers, including Hook & Cobban (1981), who note *Kamerunoceras*, *Pseudaspidoceras*, *Vascoceras* and *Fagesia* from the *juddii* Zone in the U.S., Wright & Kennedy (1981) who record *Thomasites*, *Nigericeras* and *Kamerunoceras* from below the *juddii* Zone in England and Lewy et al. (in press) who record the co-occurrence of *Metoicoceras geslinianum* (d'Orbigny, 1850) and *Vascoceras cauvini* Chudeau, 1909 in Israel.

## Coniacian

Careful reading of Coquand shows that the grounds of the Seminary at Richemont are the type locality. Fine as this outcrop is, it is singularly useless for collecting fossils. Temporary excavations at this locality in 1982 provided clean sections across the Turonian-Coniacian boundary, and showed Turonian rudistid limestones capped by a hardground and overlain by glauconitic sands of the basal Coniacian piped down into the underlying limestones in *Thalassinoides*

Ammonite Zone	Substage	Arnaud Unit
<i>Paratexanites serratomarginatus</i> <i>Gauthiericeras margae</i>	Upper Coniacian	L <sup>2</sup>
<i>Peroniceras (Peroniceras) tridorsatum</i>	Middle Coniacian	L <sup>1</sup>
<i>Forresteria (Harleites) petrocoriensis</i>	Lower Coniacian	K

Table 4. Coniacian Zones.

burrows. This discontinuity surface extends across the Aquitaine basin.

De Grossouvre recognised a 3-fold division of the Coniacian in 1894, but subsequently abandoned this, it would seem as a result of the apparent co-occurrence of species in the basal limestones of the Craie de Villedieu (which we now know to be condensed; *fide* Jarvis, Gale & Clayton 1982; Kennedy in press) and the misidentification of ammonites from the Grès Verts de Dieulefit. The 'standard' zones of the Coniacian Stage given by de Grossouvre (1901) are based on successions in Aquitaine, Touraine and Drôme, and are as follows:

*Mortoniceras emscheris* Zone (above)

*Barroisiceras haberfellneri* Zone (below)

Neither index species occurs in France. I now recognise the following standard zonation and substages shown in table 4 on the basis of outcrops in Aquitaine plus subsidiary information from Sarthe, Touraine, Dieulefit (Drôme) and Beausset (Var).

A full revision of the more than 40 ammonite species occurring in Aquitaine is given by Kennedy (in press) (see also Kennedy, Wright & Klinger 1983, and Klinger & Kennedy 1984). Table 5 summarises the known distribution of Coniacian ammonites in France as a whole.

Ammonites are locally common at the base of the sequence in Aquitaine (essentially Arnaud's Assize K), especially in the Marnes de Montignac and Gourde de l'Arche in the southeastern part of the basin. Most are *F. (H.) petrocoriensis* (Co-

quand, 1859), the '*Barroisiceras haberfellneri*' of authors from de Grossouvre (1894) to Séronie-Vivien (1972), and I know of only single specimens of *Scaphites (Scaphites)* cf. *meslei* de Grossouvre, 1894, *Metatissotia desmoulinsi* (de Grossouvre, 1894) and *M. nanclasi* (de Grossouvre, 1894). The dominance of but one ammonite species raises certain problems; while the base of the Coniacian should undoubtedly be defined, in faunal terms, by the appearance of *F. (H.) petrocoriensis*, it may be that unfavourable environmental conditions in Aquitaine excluded other species and that work elsewhere may show that a more diverse assemblage characterises the *petrocoriensis* Zone.

The base of the *P. (P.) tridorsatum* Zone is marked by the appearance of diverse *Peroniceras (Peroniceras)* species. The fauna of this zone is widely recorded in Arnaud's Assize L<sup>1</sup> in Aquitaine and is marked by *P. (P.) tridorsatum* (Schlüter, 1867) and other *Peroniceras (Peroniceras)* species with rarer *P. (Zuluiceras) bajuvaricum* (Redtenbacher, 1873). Other ammonites are *Onitschoceras? ponsianum* (de Grossouvre, 1894), *Tongoboryceras* sp. nov., *Placenticeras fritschi* de Grossouvre, 1894, *Forresteria (Harleites) nicklesi* (de Grossouvre, 1894), *Tissotia redtenbacheri* (de Grossouvre, 1894), *Metatissotia slizewiczi* (de Grossouvre, 1894), *Metatissotia ewaldi* (von Buch, 1847), *Metatissotia nodosa* (Hyatt, 1904), *Tissotioides haplophyllus* (Redtenbacher, 1873), *Scaphites meslei* (de Grossouvre, 1894), *Otoscapites arnaudi* (de Grossouvre, 1894) and nodose *Baculites*. Specimens of *Gauthiericeras? boreaui*

	petrocoriensis	tridorsatum	margae	serratomarginatus
<i>Onitshoceras ? ponsianum</i> (de Grossouvre, 1894)	-	x	-	-
<i>Tongoboryceras</i> cf. <i>tongoboryense</i> (Collignon, 1952)	-	x	-	-
<i>Tongoboryceras</i> sp. nov.	-	?	?	?
<i>Menabonites</i> sp.	-	-	-	x
<i>Pachydiscoides janeti</i> (de Grossouvre, 1894)	?	?	?	-
<i>Placenticeras fritschi</i> de Grossouvre, 1894	-	x	-	-
<i>Placenticeras semiornatum</i> (d'Orbigny, 1850)	-	-	-	x
<i>Forresteria</i> ( <i>Forresteria</i> ) <i>alluaudi</i> (Boule, Lemoine and Thévenin, 1907)	-	x	-	-
<i>Forresteria</i> ( <i>Harleites</i> ) <i>petrocoriensis</i> (Coquand, 1859).	x	-	-	-
<i>Forresteria</i> ( <i>Harleites</i> ) <i>nicklesi</i> (de Grossouvre, 1894).	-	x	-	-
<i>Peroniceras</i> ( <i>Peroniceras</i> ) <i>tridorsatum</i> (Schlüter, 1867).	-	x	-	-
<i>Peroniceras</i> ( <i>Peroniceras</i> ) <i>lepeei</i> (Fallot, 1885)	-	x	-	-
<i>Peroniceras</i> ( <i>Peroniceras</i> ) aff. <i>lepeei</i> (Fallot, 1885).	-	x	-	-
<i>Peroniceras</i> ( <i>Peroniceras</i> ) <i>subtricarinatum</i> (d'Orbigny, 1850)	-	x	-	-
<i>Peroniceras</i> ( <i>Peroniceras</i> ) <i>westphalicum</i> (Von Strombeck, 1859).	-	x	-	-
<i>Peroniceras</i> ( <i>Peroniceras</i> ) <i>dravidicum</i> (Kossmat, 1895)	-	x	-	-
<i>Peroniceras</i> ( <i>Peroniceras</i> ) sp.	-	x	-	-
<i>Peroniceras</i> ( <i>Zuluiceras</i> ) <i>bajuvaricum</i> (Redtenbacher, 1873)	-	x	-	-
<i>Peroniceras</i> ( <i>Zuluiceras</i> ) <i>isamberti</i> (Fallot, 1885)	-	x	-	-
<i>Peroniceras</i> ( <i>Zuluiceras</i> ) sp. nov.	-	x	-	-
<i>Gauthiericeras margae</i> (Schlüter, 1867)	-	-	x	-
<i>Gauthiericeras nouelianum</i> (d'Orbigny, 1850)	-	-	x	-
<i>Gauthiericeras ? boreau</i> (de Grossouvre, 1894)	-	?	x	-
<i>Protexanites</i> ( <i>Protexanites</i> ) <i>bourgeoisi</i> (d'Orbigny, 1850)	-	-	x	x
<i>Protexanites</i> ( <i>Protexanites</i> ) <i>bontanti</i> (de Grossouvre, 1894)	-	-	-	x
<i>Paratexanites zeileri</i> (de Grossouvre, 1894)	-	-	?	-
<i>Paratexanites serratomarginatus</i> (Redtenbacher, 1873)	-	-	-	x
<i>Tissotiooides</i> ( <i>Tissotiooides</i> ) <i>haplophyllus</i> (Redtenbacher, 1873)	-	x	-	-
<i>Metatissotia ewaldi</i> (Von Buch, 1878)	-	x	-	-
<i>Metatissotia slizewiczi</i> (Fallot, 1885)	-	x	-	-
<i>Metatissotia nodosa</i> Hyatt, 1903	-	x	-	-
<i>Metatissotia redtenbacheri</i> (de Grossouvre, 1894)	-	x	-	-
<i>Metatissotia desmoulini</i> (de Grossouvre, 1894)	x	-	-	-
<i>Metatissotia ? nanclasi</i> (de Grossouvre, 1894)	x	-	-	-
<i>Phlycticeroceras trinodosum</i> (Geinitz, 1850)	-	-	-	x
<i>Tridenticeras</i> sp.	-	-	x	-
<i>Eubostrychoceras</i> sp.	?	?	?	-
<i>Scalarites</i> sp.	-	x	-	-
<i>Baculites incurvatus</i> Dujardin, 1837	-	x	x	x
<i>Baculites</i> cf. <i>brevicosta</i> Schlüter, 1876	-	?	?	-
<i>Scaphites</i> ( <i>Scaphites</i> ) <i>meslei</i> (de Grossouvre, 1894)	x	x	-	-
<i>Otoscapites arnaudi</i> (de Grossouvre, 1894)	-	x	-	-

Table 5. Zonal distribution of ammonite species in the French Coniacian (Aquitaine, Dieulefit and Beausset Basins and Sarthe and Touraine)

(de Grossouvre, 1894) from L in Aquitaine may be from this zone or the succeeding *G. margae* Zone (see below).

The base of the *G. margae* Zone is marked by the appearance of the index species and *G. nouelianum* (d'Orbigny, 1850). Also present is *G. boreau*, which perhaps appears first in the underlying *P. (P.) tridorsatum* Zone. *Protexanites*

(*Protexanites*) *bourgeoisi* (d'Orbigny, 1850) first appears in the upper part of the Zone. The *margae* Zone is represented in Arnaud's Assize L<sup>2</sup>, while it also occurs in the higher parts of the bioclastic limestones of the Coniacian north-east of Saintes (e.g. St. Porchaire).

The base of the *Paratexanites serratomarginatus* Zone is marked by the appearance of the

index species. *Protexanites (Protexanites) bontanti* (de Grossouvre, 1894) is restricted to the zone, while *P. (P.) bourgeoisi* is relatively frequent. This zone is poorly represented in Aquitaine, but characteristic species occur in the higher parts of the dolomitic limestones of the Coniacian near Saujon.

The base of the Coniacian in terms of ammonites is thus the base of the classic 'habersfellneri', rightly *petrocoriensis* Zone, as is universally accepted. I have proposed elsewhere (Kennedy in press), outcrops at Antignac (Charente-Maritime), the Carrière de Chancelade near Périgueux (Dordogne), the railway cutting at St Cirq, 18 km WNW of Sarlat-la-Canéda (Dordogne), the roadcut at Aubas (Dordogne), 40 km ESE of Périgueux, Carlux (Dordogne) and Sauveterre-la-Lémace (Lot-et-Garonne) as reference sections for the zone. They are, however, unsuitable as boundary stratotypes because there is a discontinuity between Coniacian above and Turonian rudistid limestones and marls below.

A possible solution may be to reinvestigate central European sequences, notably the Priesener Schichten of Czechoslovakia and the North German outcrops in order to link precisely ammonite ranges with those of other groups that may be of more practical value in the field, e.g. *Inoceramus* species.

## Santonian

At Javrezac, a locality mentioned by Coquand as a type section for both Coniacian and Santonian Stages, the boundary between them is easily drawn at a hardground between glauconitic limestones of the Coniacian below and marls of the Santonian above.

De Grossouvre (1901) recognised a two-fold zonal division:

*Placenticerias syrtale* Zone (above)  
*Mortoniceras texanum* Zone (below)

Neither species occurs in France, while defining the base of the Santonian in ammonite terms is very difficult in Aquitaine, because the group is rare and most specimens are poorly documented. Madame M. Séronie-Vivien has kindly allowed

me to study her precisely localised specimens which, with other records, suggest that the Santonian may correspond to the range zone of *Placenticerias polyopsis* (Dujardin, 1837), which is the *P. syrtale* of authors (fide Kennedy & Wright 1983).

De Grossouvre (1894) records *Paratexanites ('Mortoniceras') serratomarginatus* (Redtenbacher, 1873) from both the Coniacian and Lower Santonian; in 1901 he decided that the Coniacian specimens all belonged to *Paratexanites ('Mortoniceras') emscheris* (Schlüter, 1876) and that *serratomarginatus* was exclusively Santonian. He never illustrated Santonian *serratomarginatus*, and none survive in museum collections that can be attributed to Arnaud's M<sup>1</sup>. Indeed, later work has shown that *serratomarginatus* and *emscheris* are not synonyms and are both from high in the Coniacian (Wiedmann 1979; Kennedy, Klinger & Summesberger 1981). De Grossouvre (1901) also indicated in a table that *Protexanites ('Mortoniceras') bourgeoisi* (d'Orbigny, 1850) extended into the lower part of the Santonian but makes no mention of this in the text and no specimens survive from M<sup>1</sup>.

Far more widely recorded is *Texanites texanus* (Roemer, 1852) and this was taken as the index of the lower zone of the Santonian by de Grossouvre. Collignon (1948) correctly interpreted Roemer's species; all European records, he maintained, belonged to two distinct varieties, regarded by subsequent workers as either subspecies of *texanus* or separate species (I believe them specifically distinct from *texanus*). Material from M<sup>1</sup> in Aquitaine (de Grossouvre 1894, p. 80, pl. 16, figs. 2, 4) was referred to the variety *gallica* Collignon, 1948; the two figured syntypes are actually from M<sup>2</sup> at Niel-le-Virouil (Charente-Maritime). The only other specimens seen from Aquitaine are two specimens labelled *Ammonites coniaciensis* Coquand, 1859 in the École des Mines Collection (now in the Collections of the Université Claude-Bernard, Lyons) ex Boucheron Collection, from La Valette (Charente), a locality referred to Assize L<sup>2</sup> by Arnaud, and a specimen in the Sorbonne Collections (now in the Université Pierre et Marie Curie, ex Toucas Collection), from M<sup>1</sup>, at Les Rentes Cognac. This specimen is a crushed *Texanites (Texanites)* sp. *Boehmoceras* occurs in M<sup>2</sup> at Nieul-le-Virouil while *P. paraplanum* (Wiedmann, 1978) and *Eu-*

*lophoceras* occur well above the base of the stage in the autoroute excavations (I thank J. M. Hancock and F. Amédéo for allowing me to see these specimens).

There thus seems no possibility of establishing an ammonite zonation of the Santonian in Aquitaine at this time.

The recent work of Bilotte, Calandra & Collignon (1971) and Bilotte & Collignon (1983) retains, to a degree, the zonal division of de Grossouvre with a Lower Santonian *Texanites texanus* Zone, a Middle Santonian Zone of *Texanites texanus*, *Placenticeras syrtale* and *Eupachydiscus isculensis* and an Upper Santonian Zone of *Placenticeras syrtale* and *Eupachydiscus isculensis*. The *Texanites 'texanus'* of the Corbières are actually *T. gallicus*, the *Placenticeras 'syrtale'* are *P. polyopsis* and *P. paraplanum*.

The ammonite distributions in Aquitaine and the Corbières are in marked disagreement, suggesting that no widely applicable zonation can be proposed on the basis of western European records at this time, pending further systematic and stratigraphic study.

In spite of these problems, there is wide agreement among ammonite workers in areas as far apart as Texas (Young 1963), Japan (Takayanagi & Matsumoto 1981), southern Africa (Kennedy & Klinger 1975, Klinger & Kennedy 1980), Madagascar (Collignon 1966) and the Middle East that the appearance of *Texanites (Texanites)* is a good indicator of the Santonian, and I would propose the appearance of this genus as the ammonite marker for the base of the stage, but hesitate to suggest a boundary stratotype (Herbert Klinger and I recorded *Texanites vanhoepeni* Klinger and Kennedy as occurring in our Coniacian V to Santonian I-II in Zululand; this misleading record indicates no more than the presence of that species in the interval mentioned and is not firm evidence for Upper Coniacian *Texanites*).

## Campanian

The problems surrounding the actual type area(s) of this stage in northern Aquitaine are a matter of some debate, and I defer to the observations of Dr. M. Neumann.

The standard ammonite zonation widely quoted is that based on the work of Schlüter (1871-1876), de Grossouvre (1894, 1901) and Haug (1911):

*Bostrychoceras polyplacum* Zone (youngest)  
*Hoplitoplacenticeras vari* Zone  
*Menabites (Delawarella) delawarensis* Zone  
*Diplacmoceras bidorsatum* Zone (oldest)

The nomenclature of the index species presents problems enough: *Diplacmoceras* (correctly *Diplacomoceras*) is a synonym of *Placenticeras* (Kennedy & Wright 1983); European *Menabites (Delawarella) delawarensis* (Morton, 1830) may not belong to Morton's species (the name *campaniensis* de Grossouvre, 1894 is available); *Ammonites vari* Schlüter, 1872 is a junior synonym of *Ammonites marroti* Coquand, 1859, while no *Placenticeras bidorsatum* or *Bostrychoceras polyplacum* have ever been figured from Aquitaine. Add to this the recent work of Blaskiewicz (1980) who defined zones of *Didymoceras donezianum* and *Nostoceras pozaryskii* above the *polyplacum* Zone in the Middle Vistula region of Poland and the observations of Hancock & Kennedy (1981), who calculated that, if the duration of the Campanian is as great as is generally accepted, then, by comparison with the older stages of the Upper Cretaceous, 20-25 ammonite zones should be recognised!

So far as the type area and elsewhere in Aquitaine are concerned, it is the collections of Arnaud that form the basis of the 'standard' zonal succession, and it was inland railway cuttings that provided much of the key material, for the beautiful coastal sections are relatively barren of ammonites.

I was able to relocate most of the Arnaud/de Grossouvre material in the collections of the Sorbonne, now housed in the Université Pierre et Marie Curie, Paris in 1981-1982; revision is still in progress, but some general observations are possible. The assemblages listed by de Grossouvre are in the correct sequence, but they do not form a zonal succession in the commonly understood sense of that term. Ammonites (other than *Baculites*) are restricted to narrow horizons, separated by barren chalk, and it is for this reason that the divisions are so coarse. The taxonomy of French Campanian ammonites is in a

STAGE	Camp.	Maastrichtian										The classic					
SUBSTAGE	Upper (part)																
FORMATION		Gulpen					Maastricht					localities					
LITHOSTRATIGRAPHIC UNIT		Vaals	Zeven Wegen	Beutenaken	Vijlen	Lixhe 1	Lixhe 2	Lixhe 3	Lanaye	Valkenburg	Gronsveld	Shiepersberg	Emael	Nekum	Meersen	Geulhem	Kunraed
SPECIES																	
<u>Saghalinites cala</u> (Forbes, 1846)																	X
<u>Pachydiscus gollevillensis</u> (D'Orbigny, 1850)																	X
<u>Pachydiscus jacquoti</u> (Seunes, 1890)																	X
<u>Pachydiscus fresvillensis</u> (Seunes, 1890)														?	?		X
<u>Hoplitoplacenticeras marroti</u> (Coquand, 1859)	X																
<u>Hoplitoplacenticeras</u> sp.	X																
<u>Sphenodiscus binkhorsti</u> (Bohm, 1898)															X	X	
<u>Baculites anceps</u> Lamarck, 1822																X	X
<u>Baculites vertebralis</u> Lamarck, 1801									X					X			X
<u>Baculites carinatus</u> Binkhorst, 1861														?	?		
<u>Baculites</u> sp. 1	X																
<u>Baculites</u> sp. 2	X																
<u>Baculites</u> sp. 3									X								
<u>Eubaculites vagina</u> (Forbes, 1846)														?	?		
' <u>Scaphites romeri</u> ' Auctorum														X			X
<u>Hoploscaphites c. constrictus</u> (J. Sowerby, 1817)					?									X		X	X
<u>Hoploscaphites c. crassus</u> (Lopuski, 1911)														X		?	
<u>Acanthoscaphites trinodosus</u> (Kner, 1850)	X																
<u>Acanthoscaphites pungens</u> (Binkhorst, 1861)																X	
<u>Acanthoscaphites verneuillanus</u> (D'Orbigny, 1842)																	X
<u>Acanthoscaphites</u> spp.																X	
<u>Diplomoceras cylindraceum</u> Auctorum														X	?		X
' <u>Glyptoceras</u> ' sp.																	X
<u>Nostoceras</u> sp.														X			

Table 6. Ammonite records from the Maastricht area.

state of chaos at present. The appearance of *Placenticerus bidorsatus* appears to be a good potential marker for the base of the stage, the species evolving paedomorphically from *P. paraplanum* of the Upper Santonian. Unfortunately the species is known from a handful of localities only (de Grossouvre 1894). The "delawarensis" Zone is generally barren, while the true identity of the rare French material and its relationship to the American species is unresolved. Ammonites are not uncommon in the *vari*, correctly *marroti* Zone, at some localities. There are, however, at least four *Hoplitoplacenticerus* species present in Aquitaine and the old records of *vari* are unreliable. The *polyplacum* Zone is well represented by new material from the environs of Royan (I thank Dr. P. Moreau for allowing me to study these specimens), and ammonites can still be collected on the coast to the south, at the top of Arnaud's P<sup>3</sup>.

Ammonites become rare in the tuffeau facies with *Orbitoides* that succeed P<sup>3</sup>. Arnaud's Assize Q is still within the Campanian, yielding typical *polyplacum* Zone forms at many localities. The highest ammonite fauna found in Aquitaine is that of the tuffeau facies around Maurens (Dordogne) and the Silex de Mussidan (Dordogne). *Baculites* are common, but are not *Baculites anceps* (Lamarck, 1799), a Maastrichtian species to which they were referred by de Grossouvre, while sphenodiscids present (*S. ubaghsi* de Grossouvre, 1894 and *S. rutoti* de Grossouvre, 1894) represent a group that occurs in both Upper Campanian and Maastrichtian according to Zaborski (1983).

Unpublished observations by Dr. M. Bilotte and myself suggest that these sphenodiscids may occur below the first appearance of *Hoploscaphites constrictus* (J. Sowerby, 1817) in the Petit Pyrénées and might thus be construed to be Upper Campanian. The *H. constrictus* include rather late forms of the species, while the sphenodiscids occur above a single *Hoploscaphites* compared to *H. pumilis* (Stephenson, 1941). In New Jersey this occurs with pachydiscids identified as *Pachydiscus* cf. *neubergicus* (Von Hauer, 1858), a species I will propose below as a possible Maastrichtian marker.

Definition of the base of the Campanian in ammonite terms presents considerable difficulties. The appearance of *Placenticerus bidor-*

*satus* (Roemer, 1841) corresponds to the 'classic' definition, but this species is so rare in the Aquitaine Basin as to render it impractical. Alternative sections might be found in the Germanies, but, even here, there are few examples, while the species appears to be restricted to north-western Europe. More widely occurring are representatives of the Texanitinae, and the evolution of *Submortonicerus* from *Texanites* is probably the most satisfactory marker in this group. *Submortonicerus* is known from Spain, Zululand, Madagascar, Mexico, the Gulf and Pacific Coasts of the United States and British Columbia, but not from the type areas of France. As a third possible group, the Scaphitidae include an important Santonian-Campanian lineage in the form of the group of *Scaphites hippocrepsis* (De Kay, 1827). This has been the subject of a masterly review by Cobban (1969). Specimens referred to as *Scaphites aquisgranensis* (Schlüter, 1872) by European authors that occur with *Placenticerus bidorsatus* in Europe are regarded by Cobban as representatives of the latest of three distinct forms of *S. hippocrepsis* (*S. hippocrepsis* III), which occurs widely in the U.S. Western Interior, the Atlantic and Gulf Coast plains, the Aquitaine Basin, West Germany, Belgium and Holland.

## Maastrichtian

The type section of the Maastrichtian Stage in St. Pietersberg near Maastricht in eastern Holland, where only the upper part of the stage is present, and even the uppermost Maastrichtian is absent. A type section has been designated at the ENCI quarry. The lower boundary of the Maastrichtian, where exposed in the Maastricht area, is a marked discontinuity surface. As a result, Surlyk (1975, 1982) and Schulz (1978) proposed a boundary stratotype at Kronsmoor in north Germany, corresponding, at that locality, to the base of the *Belemnella lanceolata* Zone, with the first *Hoploscaphites constrictus* appearing 3.5 to 5 m above the base.

Standard zonations of the Maastrichtian have not been widely accepted. One definition has been to consider the stage to correspond to the total range zone of *Hoploscaphites constrictus*; Wright (in Arkell et al. 1957) gave a zonation of:

*Sphenodiscus* sp. (above)

*Pachydiscus neubergicus* (below)

for the "Cretaceous of Classic Areas of Western Europe." Given a Maastrichtian that lasted around 7 Ma (Odin 1983) and the presence of genera (*Hoploscaphites*, *Baculites*) that provide fine biozonations elsewhere in the world, the situation is an unsatisfactory one. So far as the area around Maastricht is concerned my observations are of a preliminary nature only; table 6 summarises the ammonite records gleaned from published literature and the old collections at Brussels and Maastricht.

Defining the base of the Maastrichtian by ammonites has several possible solutions. The Kronsmoor section in the White Chalk facies of north Germany has obvious attractions as a western European boundary stratotype, taking the base of the *Belemnella lanceolata* Zone as the base of the stage which can in turn be correlated with the first occurrence, a few metres higher, of *Hoploscaphites constrictus* as well as with echinoid and benthic foram zones. It must be remembered, however, that this Boreal sequence of faunas cannot easily be applied to the greater part of the world. Should we instead, at least consider a group that occurs in the European White Chalk region, in the north Tethys and indeed, in the southern hemisphere, that is to say the ammonite family Pachydiscidae? Specifically, should we take the first appearance of *Pachydiscus neubergicus* (von Hauer, 1858) as the ammonite indicator of the base of the Maastrichtian? The advantages of this solution are:

1. It occurs in the European White Chalk sequence (Denmark, N. Germany) and can be linked to the belemnite sequences there.
2. It occurs in the tectonically complex areas of Alpine Europe (the type locality, Neuberg, Styria, associated with the widespread Maastrichtian marker *Eubaculites* plus *Diplomoceras cylindraceum* (d'Orbigny, 1842) and *Hoploscaphites tenuistriatus* (Kner, 1848)).
3. It occurs in north Tethyan pelagic facies with good *Inoceramus*, micro- and nannofossil associations (Zumaya, N. Spain; Landes, France).
4. It occurs in eastern North America associated with widely recorded *Nostoceras* such as *N. hyatti* Stephenson, 1941.
5. It occurs in the Indo-Pacific region associ-

ated with diverse ammonites and other groups (southern India, Zululand, South Africa).

6. In Zululand, South Africa it occurs with diagnostic elements of the New Zealand – W. Australia – Antarctica 'Upper Campanian' *Gunnarites* fauna.

As a boundary section the Zumaya sequence may prove a satisfactory one, given the extensive faunas now being studied by Drs. Ward and Wiedmann. It must be stressed, however, that *P. neubergicus* may have a long vertical range in the Maastrichtian, for it is a species that spans the *Belemnella occidentalis* / *Belemnitella junior* Zone boundary in Denmark (Birkelund 1979), and that mere occurrence is not an indication of the base of the stage.

## Conclusions

In ammonite terms, the possibility that the bases of the Upper Cretaceous stages can be satisfactorily defined as follows is proposed for further study and discussion:

1. Cenomanian: appearance of either *Hypoturrites* or *Neostlingoceras* lineage in the Tunisian sections.
2. Turonian: appearance of *Pseudaspidoceras flexuosum* in the west Texas – New Mexico area.
3. Coniacian: appearance of *Forresteria* (*Harleites*) *petrocoriensis* in N. German/Czech sequences.
4. Santonian: appearance of *Texanites*.
5. Campanian: appearance of *Placentoceras bidorsatum* in Europe or *Submortonoceras* elsewhere or on the basis of evolutionary changes within the *Scaphites hippocrepis* lineage.
6. Maastrichtian: appearance of *Pachydiscus neubergicus*.

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## Dansk sammendrag

Brugen af ammonit »standard zonerne« i øvre kridt etagerne i Vesturopa vurderes kritisk. Det foreslås at definere øvre kridts etager på grundlag af første optræden af følgende ammoniter: cenomanien: *Hypoturrillites* eller *Neostlingoceras*; turonien: *Pseudaspidoceras flexuosum*; Coniacien: *Forresteria petrocoriensis*; santonien: *Texanites*; campanien: *Placentoceras bidorsatum* eller *Submortonoceras* eller stadier af *Scaphites hippocrepis* udviklingslinjen; maastrichtien: *Pachydiscus neubergicus*.

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