

Quaternary palaeoceanography and palaeogeography in northern Denmark: a review of results from the Skagen cores

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Knudsen, K. L., Conradsen, K., Heier-Nielsen, S. & Seidenkrantz, M.-S.:

Quaternary palaeoceanography and palaeogeography in northern Denmark: a review of results from the Skagen cores. *Bulletin of the Geological Society of Denmark*, Vol. 43, pp. 22–31. Copenhagen, 1996–07–14.
<https://doi.org/10.37570/bgsc-1996-43-03>

Palaeoenvironmental reconstructions from the Skagen record contribute to the understanding of Late Quaternary climatic changes and variations in the oceanographic circulation pattern in the entire North Atlantic region. The Skagen cores penetrated c. 192 m of Quaternary sediments comprising two marine Late Quaternary records: A 7 m marine unit (185.3–178.3 m) comprised the entire last inter-glacial, including its lower and upper transitions (Late Saalian-Eemian-Early Weichselian), while the upper 132 m of marine deposits covered the last about 15,000 years from the Late Weichselian through the Holocene, including the Pleistocene-Holocene transition. Results from the study of lithology, foraminifera, stable isotope measurements and radiocarbon dates are reviewed while emphasizing the most important contributions to the general understanding of the North Atlantic Quaternary history.

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Introduction

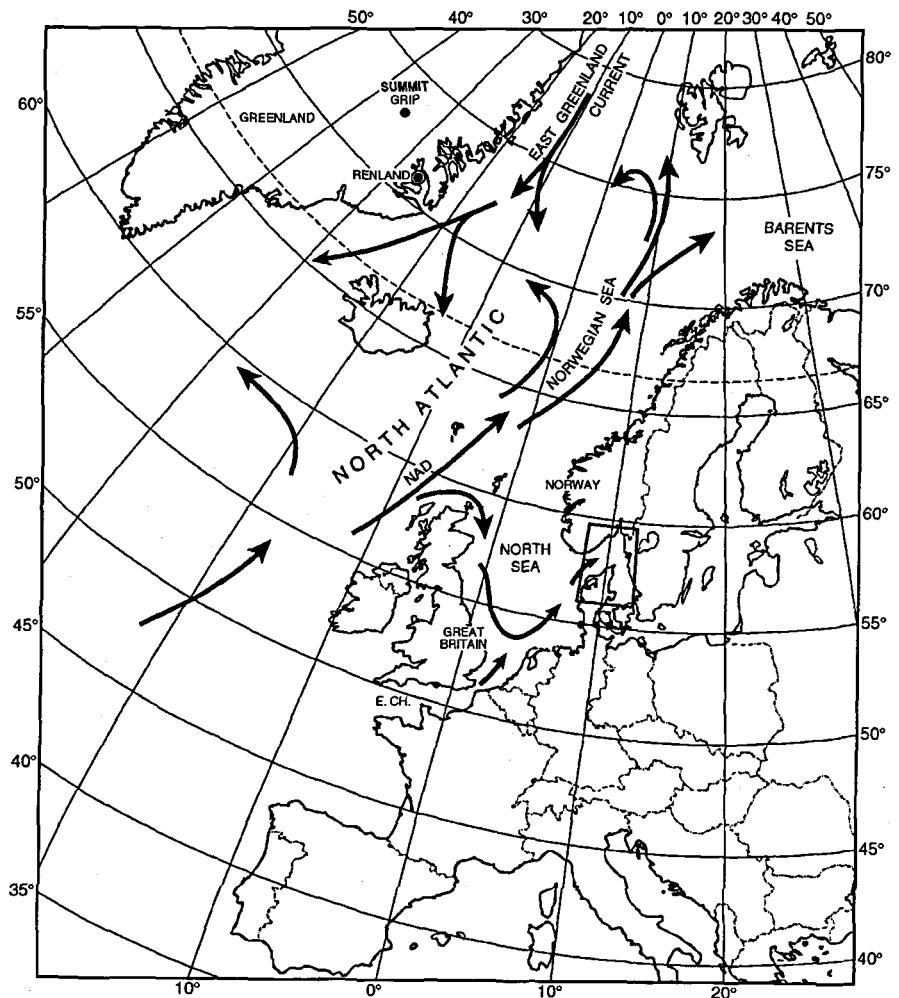
The North Atlantic current system plays an important role in the understanding of both global and regional climatic systems of the past (e.g. Lehman & Keigwin 1992; Bond et al. 1993; Weaver & Hughes 1994). As changes in northwest European climate and oceanography are strongly influenced by variations in the influx of Atlantic water into the Norwegian Sea and the North Sea, the NW European shelf areas appear to be key areas in the understanding of these fluctuations.

The present surface current system in the North Atlantic (Fig. 1) has probably been the prevailing oceanographic setting, not only through the Holocene, but also through at least a major part of each of the previous interglacial periods. During glacial periods, however, the Polar Front was displaced to an area off southern Europe, while the North Atlantic Current was considerably weakened, perhaps even completely 'shut down' during certain periods (e.g. Weaver & Hughes 1994; Bond 1995). These migrations of the water masses are also clearly reflected in the Danish shelf sediments.

During the Late Quaternary an up to about 300 m deep palaeo-embayment, the Kattegat Depression, stretched from the Skagerrak across northern Jutland and into the Kattegat sea (Fig. 2) (see the base Quaternary map in Lykke-Andersen 1987 and Lykke-Andersen et al. 1993a). This embayment was connected to the North Sea and the North Atlantic to the northwest and may be considered to have been a palaeo-extension of the Norwegian Trench. The palaeo-embayment has gradually been filled with sediments from the Late Saalian through to the Holocene.

This scenario resulted in the deposition of thick marine sedimentary sequences, which enable high-resolution studies of the palaeoenvironment of the late Quaternary. Previous studies of the foraminifera from the Kattegat Depression have mainly focused on the Eemian and Early Weichselian (Bahnsen et al. 1974; Knudsen & Lykke-Andersen 1982; Lykke-Andersen, A.-L. 1987; Knudsen 1992; Seidenkrantz 1993a, 1993b) and the Late Weichselian (Feyling-Hanssen et al. 1971; Knudsen 1978; Knudsen & Nordberg 1987; Bergsten & Nordberg 1993; Seidenkrantz & Knudsen 1993). In addition, Holocene deposits from the Kattegat and Skagerrak regions have previously been studied,

Fig. 1. Location of sites mentioned in the text and the present surface current system in the North Atlantic. NAD = North Atlantic Drift; E. CH. = English Channel. The area covered by Figure 2 is indicated.



e.g. Fält (1982), Nordberg & Bergsten (1988), Nordberg (1989) and Conradsen (1995). It has, however, hitherto not been possible to compare the boundary conditions of both the Eemian and the Holocene at the same site.

The purpose of this paper is to present a review of the Quaternary stratigraphy of the Skagen 3 and Skagen 4 cores (Figs 2–3) based on lithology, foraminiferal analyses, stable isotope results and AMS radiocarbon dates. In addition, we have placed emphasis on a comparison of the Eemian and Holocene records as well as the glacial-interglacial transitions at this specific site.

A short compilation of ongoing work including some preliminary results from the Skagen cores is given by Knudsen (1994). More detailed studies of the foraminiferal and stable isotope stratigraphy from the glacial-interglacial-glacial cycle at Skagen are presented by Seidenkrantz et al. (1995) and Seidenkrantz & Knudsen (unpublished). Results of foraminiferal analyses, stable isotope results and ^{14}C dates from the Late

Weichselian and Holocene sequence, including the Pleistocene-Holocene transition, are given by Knudsen et al. (1996), Conradsen & Heier-Nielsen (1995) and Heier-Nielsen et al. (1995), while geotechnical properties from the boreholes are discussed by Thorsen (1995a, 1995b) and Thorsen & Mortensen (1995).

Further palaeoenvironmental studies of the Skagen record are in progress at a number of different institutions, these being the Geological Survey of Denmark and Greenland (sediments and molluscs), University of Lund (diatoms), University College London (ostracods), University of Cambridge (pollen) and Université de Québec à Montréal (dinoflagellates). The chronology of the record beyond the range of the ^{14}C dating method is being determined at the Bergen Amino acid Laboratory and at the Luminescence Laboratory at Risø (preliminary results in Poolton et al. in press).

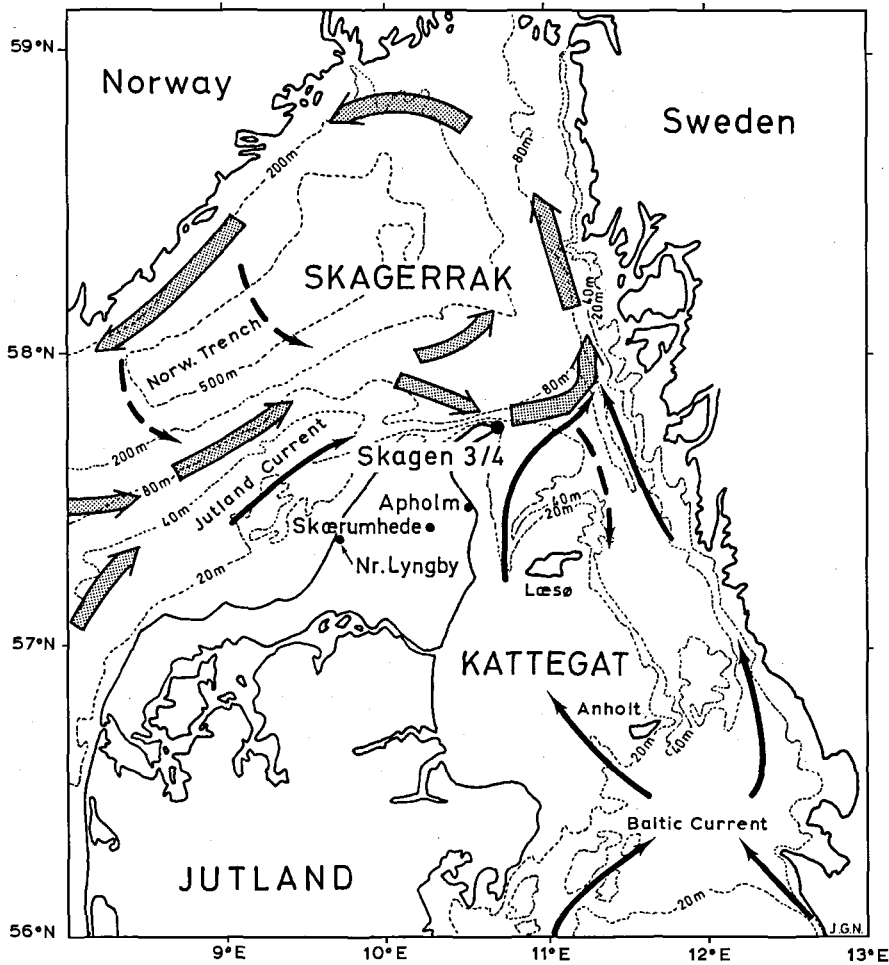


Fig. 2. Present regional oceanographic circulation pattern and location of sites mentioned in the text. Redrawn from Svansson (1975) and Nordberg (1991).

The Skagen 3 and Skagen 4 boreholes

The borings were carried out in 1992 and 1993 just north of the city of Skagen (Fig. 2) at 1 m above sea level. Two oil exploration wells had previously been drilled in the area, but the new borings are the first to provide material of the standard required to enable a detailed study of the Quaternary strata. Skagen 3 covers the interval between 30 and 220 m depth while Skagen 4 supplements with the uppermost 30 m. The Skagen 3 core reached the pre-Quaternary (Lower Cretaceous non-marine sands) at c. 192 m depth (see Knudsen 1994).

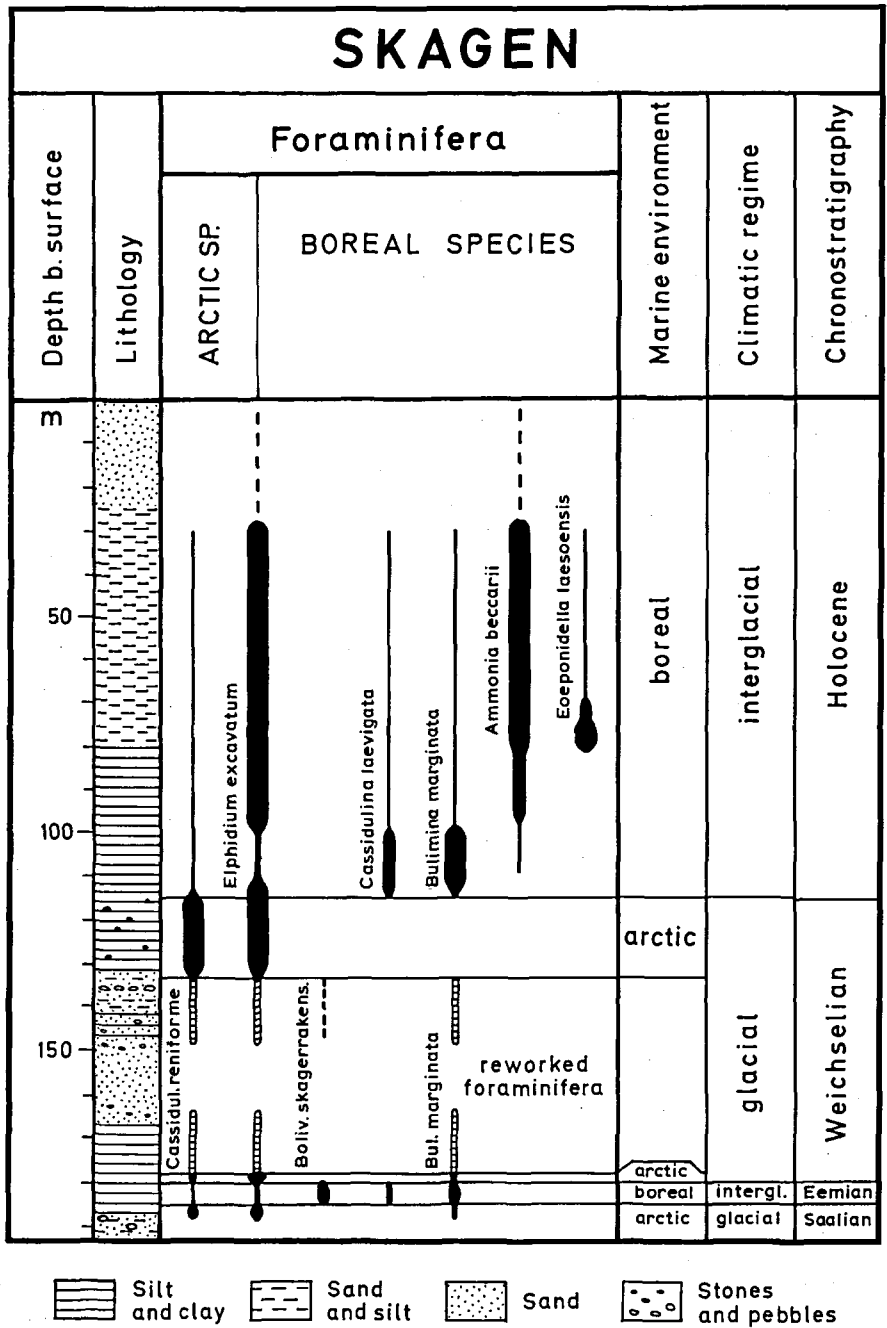
The drilling and laboratory techniques have been described by Conradsen & Heier-Nielsen (1995), Knudsen et al. (1996) and Seidenkrantz & Knudsen (unpublished).

The Saalian-Eemian-Early Weichselian cycle (192–178.3 m)

The oldest Quaternary deposits at Skagen represent non-marine glacial sands, clays and diamicts (Fig. 3), corresponding to the Saalian Glacial (oxygen isotope stage 6) (see also Knudsen 1994).

A marine transgression subsequently occurred in the Late Saalian, and marine clays with subarctic benthic foraminiferal faunas dominated by *Elphidium excavatum* (Terquem) forma *clavata* Cushman and *Cassidulina reniforme* Nørvang were deposited (Zone M, Figs 4–5).

The Saalian/Eemian transition (stage 6/5e) is marked by the gradual immigration of warm water species; first *Bulimina aculeata* d'Orbigny and *B. marginata* d'Orbigny (grouped as *B. marginata* in Figs 4–5) together with *Cassidulina laevigata* d'Orbigny (Zone L) and later, in addition, *Melonis barleeanus* (Williamson), *Bolivina skagerrakensis* (Qvale & Nigam) and



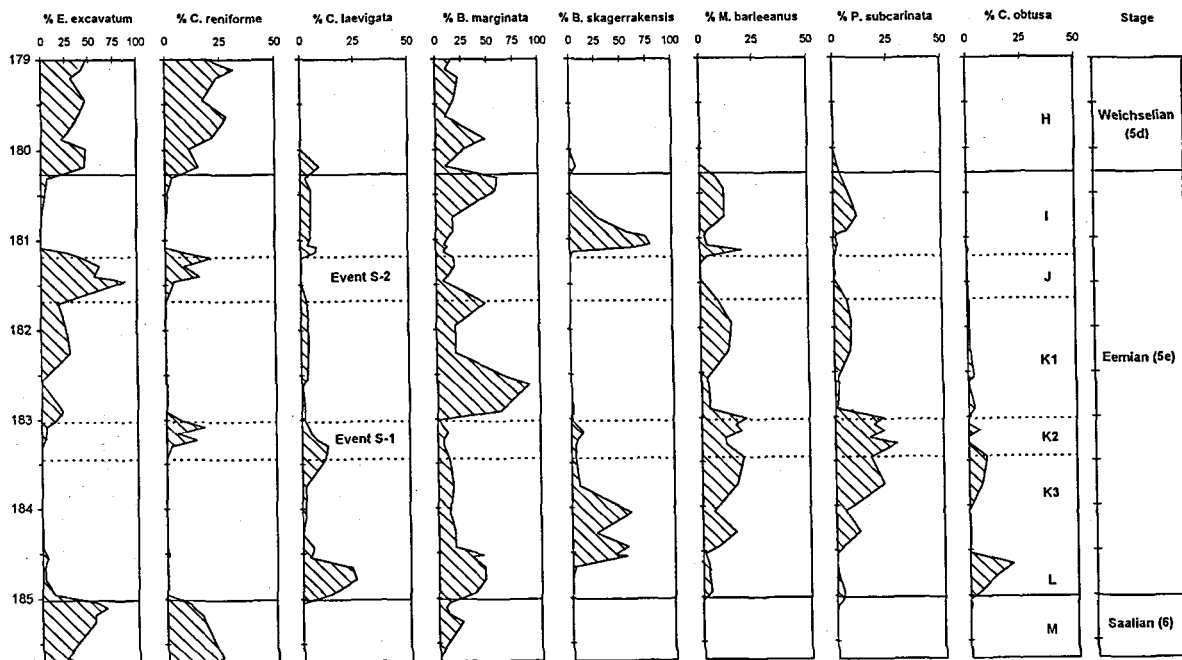


Fig. 4. Percentages of selected foraminiferal species in the glacial-interglacial-glacial cycle (Saalian-Eemian-Early Weichselian). The interval between 185.7 and 179.0 m depth in the Skagen 3 core is shown. The two mid-Eemian cooling events (Event S-1 and S-2), the foraminiferal zonation (Zones M to H) and the chronostratigraphy are indicated in the diagram. Further details in Seidenkrantz et al. (1995) and Seidenkrantz & Knudsen (unpublished).

Pullenia subcarinata (d'Orbigny) (Zone K3, Figs 4–5). These species are among the characteristic warm water elements in the main part of the 5 m thick Eemian Interglacial deposits (Zones L–I).

In two intervals within the Eemian, however, the foraminiferal assemblages differ from the remaining Eemian strata. The lower interval (Event S-1, Subzone K2), is marked by a peak in the arctic species *C. reniforme* in an otherwise temperate fauna, and the upper interval (Event S-2, Zone J) is characterized by the truly subarctic fauna of *C. reniforme* and *E. excavatum* (Fig. 4). The two cooler intervals also show excursions in stable isotope values (Seidenkrantz & Knudsen, unpublished).

These two periods of cooling are presumably the most remarkable features of the Eemian Interglacial record in the Skagen 3 core (further discussed by Seidenkrantz et al. 1995 and Seidenkrantz & Knudsen unpublished). Marine Eemian sediments from the region have been well documented for many years, e.g. from Skærumhede, Apholm, Nørre Lyngby and Anholt (Fig. 2) (Knudsen & Lykke-Andersen 1982; Knudsen 1984; Lykke-Andersen, A.-L. 1987; Seidenkrantz 1993a, 1993b), but similar coolings have not previously been recorded in the interglacial record. We presume that this is due to lack of recovery from the relevant levels in these older cores.

The Eemian sediments at Skagen are succeeded by 2 m of marine clays with a subarctic foraminiferal fauna

dominated by *E. excavatum*, *C. reniforme*, *Epistominella takayanagii* Iwasa, *B. marginata* and *B. aculeata*. This interval is allocated to part of the Early Weichselian (substage 5d, Seidenkrantz et al. 1995; Seidenkrantz & Knudsen unpublished). A more complete Lower Weichselian sequence has previously been recorded at Apholm (Knudsen 1984, 1992), but insufficient sample resolution has prevented a detailed study of the record. A study of climatic oscillations in an Early Weichselian marine record at Nørre Lyngby is, however, in progress (Kristensen et al., unpublished).

The Early to Middle Weichselian (178.3–132 m)

About 45 m of clays and sands with scattered pebbles and stones superpose the marine Lower Weichselian at Skagen. It contains a relatively badly preserved benthic foraminiferal assemblage. The fauna is dominated by a mixture of arctic species (*E. excavatum*, *C. reniforme*, *Islandiella helenae* Feyling-Hanssen & Buzas, *Haynesina orbiculare* (Brady) and *Elphidium bartlettii* Cushman) and warm water species (e.g. *B. skagerrakensis*, *B. marginata* and *C. laevigata*), suggesting that a significant number of the specimens have been reworked. The warm water species were possibly transported into an arctic, glaciomarine environ-

ment with an existing fauna. A similar 30–40 m thick basin fill containing mixed, partly mixed or perhaps fully reworked foraminiferal assemblages have previously been identified in several borings in the region, i.e. Skærumhede, Læsø and Anholt (Fig. 2) (Bahnsen et al. 1974; Knudsen 1986; Seidenkrantz 1993b), and this sequence can also be followed on seismic profiles over most of the Kattegat Depression (Lykke-Andersen et al. 1993b).

The Main Weichselian Glaciation is presumed to be indicated by a marked concentration of pebbles and stones in the uppermost part of this basin fill sediment (c. 137–135 m depth; Fig. 3).

The Late Weichselian and the Holocene (132–0 m)

A marine environment was re-established in the area at about 15,000–14,500 BP (from 132 m depth in the core). The sedimentation rate was high, and the marine clays initially contain a scatter of ice-rafted stones and foraminiferal assemblages dominated by ice-proximal species such as *E. excavatum* f. *clavata* and *C. reniforme* (Zone F; Figs 3 and 6).

The sedimentation rates decreased markedly between about 13,100 and 9,600 BP (between 117 and 114 m depth), and there was a temporary maximum of accessory species such as *Nonionellina labradorica* (Dawson) (see Fig. 6 and Knudsen et al. 1996), *Astronionion gallowayi* Loeblich & Tappan and *Cibicides lobatulus* (Walker & Jacob) in the assemblages. In addition, two other accessory species, *Stainforthia loeblichii* (Feyling-Hanssen) and *Pullenia osloensis* Feyling-Hanssen, appeared in the lower and upper part of this interval (see Knudsen et al. 1996).

The fluctuations in the foraminiferal faunas within this period, associated with fluctuations in stable isotope values (Fig. 6), show that the climatic variations of the Bølling Interstadial, the Older Dryas Stadial, the Allerød Interstadial, the Younger Dryas Stadial and the Preboreal affected the oceanographic conditions in the Kattegat-Skagerrak area (see also Knudsen et al. 1996). The bottom waters were still influenced by cold, normal saline Atlantic water masses imported through the Norwegian Trench, presumably as a compensation current to an increased outflow of meltwater from the Baltic area during the Bølling-Allerød Interstadial complex and the Preboreal. Minor changes during the Older Dryas and Younger Dryas Stadial periods in both the foraminiferal fauna and the isotopic composition indicate short periods when there was reduced stratification of the water column. Reduced meltwater outflow seems to have resulted in enhanced mixing of the water.

Thus, arctic/subarctic conditions prevailed in the deep Skagerrak during the entire Bølling-Allerød Interstadial complex, the Younger Dryas and the early part of the Preboreal. Previous studies (e.g. Knudsen

1978; Bergsten & Nordberg 1992; Conradsen 1995) have shown, however, that during the same period of time warmer foraminiferal assemblages occurred at least periodically in shallower water environments in the Skagerrak-Kattegat region.

Only minor changes occurred in the foraminiferal assemblages at the Pleistocene-Holocene transition (Subzone E3/E2 boundary; 115 m depth). The first boreal species, *C. laevigata*, had already appeared by Subzone E2 and increased through Subzone E1 together with *P. osloensis*, while *C. reniforme* and *N. labradorica* decreased (Fig. 6). In general, however, arctic/subarctic environments still prevailed in the area until about 9,600 BP.

After 9,600 BP (i.e. during the remaining part of the Preboreal and during the Boreal period; Zones D and C; Fig. 6) there was a consistent occurrence of *C. laevigata* and a pronounced increase in other boreal species such as *B. marginata* concurrent with a decrease in both *E. excavatum* and *C. reniforme*. This indicates increased bottom temperatures, reflecting a major inflow of warm Atlantic water masses into this relatively deep area of the eastern North Sea basin. Thus, full-interglacial hydrographic conditions were first established in the entire region after 9,600 BP. A similar two-stage post-Younger Dryas oceanographic transition with a first change prior to 10,000 BP and a second a few hundred years later has been previously recorded in the Skagerrak, the Norwegian Sea and the North Atlantic (e.g. Bergsten 1994; Koç Karpuz & Jansen 1992; Lehman & Keigwin 1992; Sarnthein et al. 1994).

At about 7,600 BP a distinct change occurred in the Holocene fauna (at 100 m depth in core). *B. marginata* and *C. laevigata* declined markedly (Fig. 3) and *E. excavatum* forma *selseyensis* (Heron-Allen & Earland) and *Ammonia beccarii* (Hofker) became the most significant species. This type of fauna is identical to that found in recent assemblages around Skagen (Conradsen et al. 1994), indicating the establishment of the modern circulation pattern in the North Sea and Skagerrak-Kattegat areas after opening of the English Channel (e.g. Jelgersma 1979; Lehman et al. 1991; Nordberg 1991) and the Danish Straits (e.g. Winn et al. 1986; Björck 1995). After 7,600 BP the sedimentary conditions were significantly influenced by sediment transport in the Jutland Current (including the coastal current), which caused the development of the Skagen Spit. The Holocene development of the Skagen Spit is shown in Figure 7.

A faunal change and a lithological change at about 5,500 BP (c. 81 m depth in core; Fig 3) is indicated by the immigration of a new species (*Eoepionidella laesoensis* Michelsen) into the area and lithologically by the deposition of relatively coarser sediments (see also Conradsen & Heier-Nielsen 1995). The change to higher energy conditions probably reflects a more pronounced Jutland Current and thus increased inflow of North Sea water into the Kattegat (see also e.g. Nord-

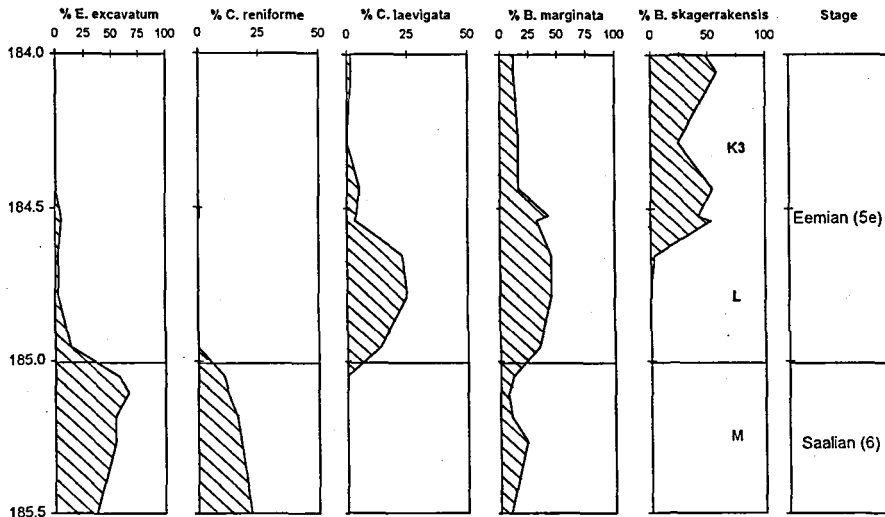


Fig. 5. Details of the distribution of selected species at the glacial-interglacial transition from the Late Saalian to the Early Eemian at around 185 m depth in the Skagen 3 core. Foraminiferal assemblage zones (Zones M to K3) and chronostratigraphy are indicated (see Seidenkrantz et al. 1995 and Seidenkrantz & Knudsen, unpublished).

berg 1991). This modification of the hydrographic system in the Skagerrak–Kattegat region is most likely a response to the climatic cooling at the end of the Holocene climatic optimum in Atlantic time and may thus reflect a regional cooling of the surface water masses of the North Atlantic.

Comparison of the Eemian and Holocene Interglacials

During the last interglacial (Eemian) the temperatures were generally a few degrees higher than those experienced during the Holocene and today in the North Atlantic region (e.g. Sejrup & Larsen 1991; Knudsen 1992; Larsen et al. 1995; Johnsen et al. 1995a). This is also confirmed by our data from Skagen (Seidenkrantz & Knudsen, unpublished). In contrast to the relatively stable temperature conditions during the Holocene, studies of the GRIP ice core from the Summit of the Greenland ice cap suggest that the Eemian Interglacial comprised significant climatic fluctuations (e.g. Dansgaard et al. 1993; GRIP Members 1993; Larsen et al. 1995; Johnsen et al. 1995b). Similar climatic variations have also been documented both in the terrestrial environment in NW Europe (e.g. Tzedakis et al. 1994; Thouveny et al. 1994), in marine deposits in the North Atlantic (e.g. Seidenkrantz & Rahman, unpublished; Fronval et al., unpublished; Sejrup et al. in press), and in NW European shelf deposits, as demonstrated by Seidenkrantz et al. (1995), Seidenkrantz & Knudsen (unpublished) and in the present paper. The two mid-Eemian cool periods may have been caused by a temporary weakening of the North Atlantic Drift (Seidenkrantz et al. 1995). In this context, a discussion of the climatic influence of east-west oscillations

of the oceanic front systems between Polar and Atlantic surface waters in the Norwegian Sea was recently presented by Sejrup et al. (in press).

In European terrestrial records, the Eemian climatic optimum seems to have occurred within the first 3000 years of the interglacial, and a similar early climatic optimum for this period has been demonstrated in the Norwegian Sea (see Larsen et al. 1995). This corresponds to the longest period with temperatures higher than at present in the GRIP ice core. A comparable early Holocene climatic optimum is known from the North Atlantic region, including the Renland ice core from East Greenland (e.g. Iversen 1973; Nesje & Kvamme 1991; Johnsen & Dansgaard 1992) and has recently also been demonstrated in the GRIP ice core from Greenland (Johnsen et al. 1995a). It has not been possible to identify a climatic optimum in the initial part of the Eemian record at Skagen in the present study.

Larsen et al. (1995) calculated that a transgression maximum occurred about 2000 years later than the Eemian climatic optimum. This has not been detectable in the Skagen record either. The general palaeo-water depth at the site was presumably too great (200–300 m) to detect any minor fluctuations in sea level through the interglacial period (Seidenkrantz & Knudsen, unpublished).

Climatic oscillations, as reflected in both the foraminiferal assemblages and the stable isotope compositions of the Eemian record, are not found in the Holocene record from Skagen. The faunal succession in the Holocene sequence does, however, demonstrate changes in the hydrographic circulation pattern at about 7,600 BP and at 5,500 BP. It is suggested that the earlier change was related to the opening of the English Channel and the Danish Straits, while the latter may reflect the general cooling in the North Atlantic region at end of the Holocene climatic optimum. The

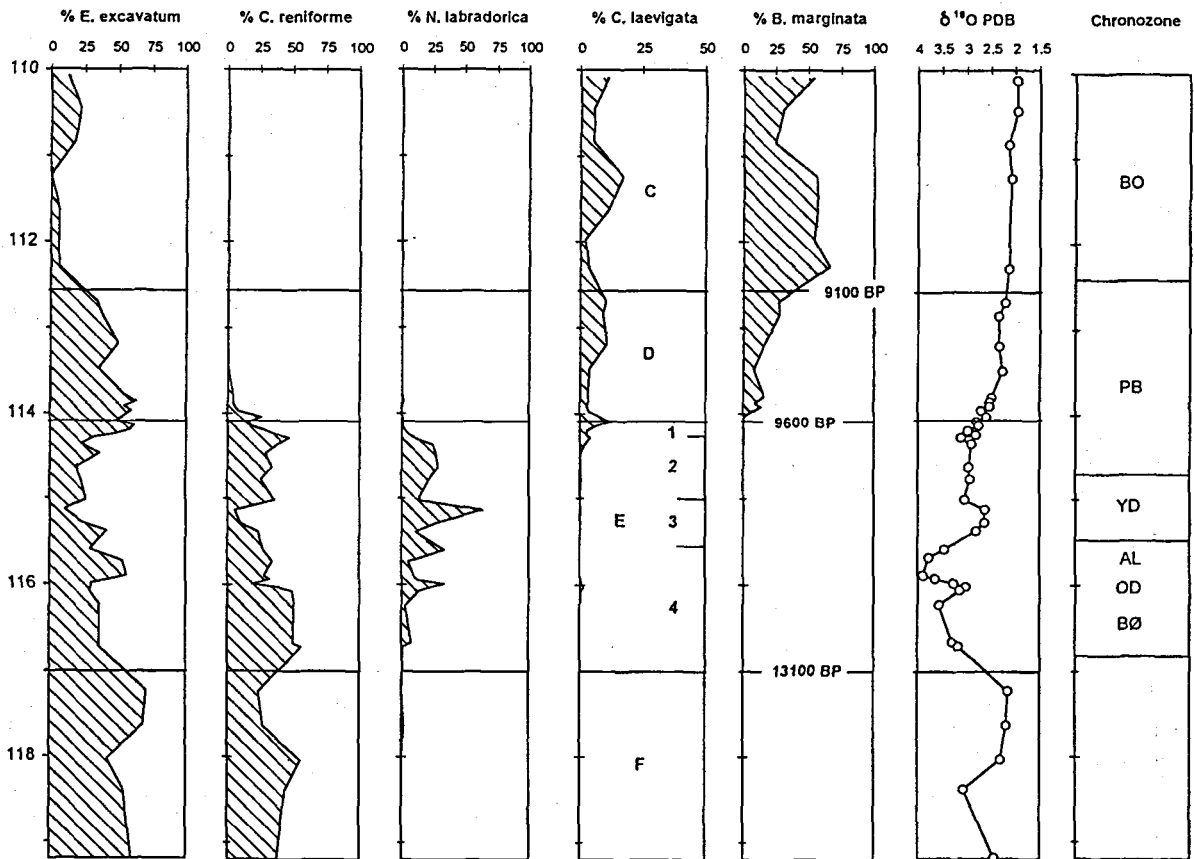


Fig. 6. Percentage distribution of selected foraminiferal species and the oxygen isotope values at the glacial-interglacial transition from the Late Weichselian to Early Holocene between 119 and 110 m depth in the Skagen 3 core. Foraminiferal assemblage zones (Zones F to C) and the age of some important zone boundaries are indicated in addition to the chronostratigraphy (further details concerning the age model, see Knudsen et al. 1996). BØ = Bølling; OD = Older Dryas; AL = Allerød; YD = Younger Dryas; PB = Preboreal; BO = Boreal.

changes in palaeo-water depth through the Holocene at Skagen were affected by both isostasy and global sea level, but the gradual decrease in water depth since about 9,000 BP is largely attributed to the extremely high sediment accumulation rates in the area (Conradsen & Heier-Nielsen 1995).

The glacial-interglacial transitions

Studies of the Late Saalian-Eemian transition in different parts of the world show that the overall gradual warming was interrupted by a short-term climatic deterioration comparable with the Younger Dryas cold spell at the Late Weichselian-Holocene transition (Seidenkrantz et al. in press). In the Danish area this has been demonstrated by the foraminiferal and stable isotope data from the Anholt core (Seidenkrantz 1993a). A similar complete record of the Late Saalian

was not found at Skagen, even though the transition from the glacial into the interglacial is present in the sequence.

In general, the faunal change at the Saalian-Eemian transition in the Skagen core (Fig. 5) is very similar to that found at the Weichselian-Holocene transition (Fig. 6). The arctic species *E. excavatum* forma *clavata* and *C. reniforme* decrease markedly in frequency, while the boreal *C. laevigata* and *B. marginata* increase during the warming. In the Eemian sequence this faunal change is succeeded by immigration of a number of deep water species, e.g. *B. skagerrakensis*, which never migrated into the area at Skagen in the Holocene because the water depth in area was too shallow at that time (Conradsen & Heier-Nielsen 1995; Conradsen et al. 1994).

The faunal changes at glacial-interglacial transitions, thus, generally seem to be much alike from one interglacial to another if the environmental conditions are comparable. Minor details in the assemblages may,

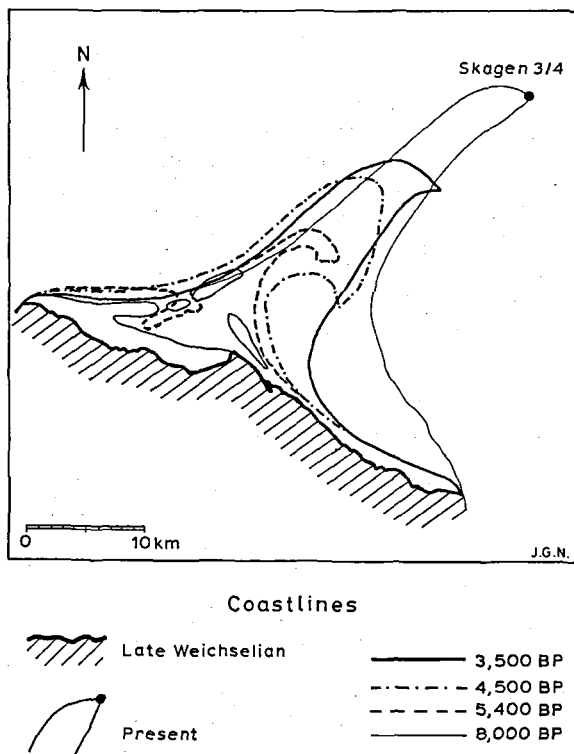


Fig. 7. Development of the Skagen Spit after about 8,000 BP. Redrawn from Hauerbach (1992).

however, prove to be important in the separation of different glacial-interglacial transitions in a particular area. Thus, the immigration of *Cassidulina obtusa* Williamson in the initial part of the Eemian Interglacial, as seen in the Skagen record (Fig. 4), might prove to be indicative of the Eemian in the region. A similar peak in *C. obtusa* is also seen at the Saalian-Eemian transition in the Apholm, the Anholt and the Nørre Lyngby cores (Knudsen 1984; Seidenkrantz 1993a; Kristensen et al., unpublished). Details in the immigration patterns at glacial-interglacial transitions are, however, not yet fully understood.

Summary of results from the Skagen cores

The high-resolution studies of the marine Quaternary sequences at Skagen contribute to the understanding of the Late Quaternary climatic and oceanographic history of the entire North Atlantic region. A summary of some of the most important observations and conclusions resulting from the study of lithology, foraminifera, stable isotopes and radiocarbon dates includes:

- Climatic oscillations in the Eemian Interglacial shelf

record comparable to those described from the GRIP ice core in Greenland and in the terrestrial record of NW Europe. These oscillations are suggested to be associated with fluctuations in the North Atlantic Drift.

- A period in which glaciomarine ice-proximal conditions prevailed in the area between about 15,000 and 13,100 BP.
- The inflow of cold, normal saline Atlantic water masses at the bottom of the Skagerrak-Kattegat region during the Bølling-Allerød, the Younger Dryas and the early part of the Preboreal (c. 13,100-9,600 BP), while higher temperature conditions prevailed in shallower areas during most of this time period. Indications of decreased stratification of the water column during the Older and the Younger Dryas cold events.
- Indication of the inflow of warm Atlantic water masses in the entire region at about 9,600 BP.
- The establishment of the present hydrographic circulation system in the North Sea and the Skagerrak-Kattegat area at approximately 7,600 BP. This is suggested to be associated with the opening of the English Channel and the Danish Straits.
- Modification of the current system at approximately 5,500 BP, which may be a result of the general cooling in the North Atlantic at the end of the Holocene climatic optimum.
- The bottom water temperatures at Skagen were generally a few degrees higher during the Eemian Interglacial than during the Holocene and at present.
- A general similarity in the faunal change at the Saalian-Eemian and the Weichselian-Holocene transitions, but with specific characteristics in the immigration pattern which might show to be important for age determinations.

Acknowledgements

The drilling at Skagen was part of a joint venture between the GeoKat project and the Skagen Spit Project and was financed by the Danish Natural Science Research Council and the Municipality of Skagen. Additional funding for the study was provided by the Carlsberg Foundation, Denmark and by the European Union SHELF-project. Our sincere thanks to Peter Kristensen for help with the laboratory work and for valuable discussions of the results, to Gillian Scott for correcting the English and to Jette Gissel Nielsen for help with the diagrams.

Dansk sammendrag

Palæomiljø-undersøgelser på grundlag af lithologi, foraminiferer, stabile isotoper og ^{14}C dateringer fra Skagen har vist sig at bidrage væsentligt til forståelsen

af ændringer i klima og strømsystemer i hele det nordatlantiske område.

Skagen ligger i den centrale del af en dyb sydøst-nordvest orienteret bassinstruktur i den præ-Kvartære overflade. Denne sydøstlige udløber af Norske Rende blev transgrederet af havet i Sen Saale og har siden været præget af periodevist meget høje sedimentationshastigheder i et marint miljø. De tykke marine lagserier i området giver grundlag for at foretage palæomiljø-undersøgelser med høj tidsopløsning.

Den Kvartære lagserie ved Skagen er ca. 192 m tyk og indeholder to bemærkelsesværdige marine enheder. I den dybere del af boringen (185,3–178,3 m) forekommer en 7 m tyk marin ler, som overlejrer glaciogene sedimenter. Denne marine lagserie repræsenterer en hel glacial-interglacial-glacial cyklus fra Sen Saale, Eem og Tidlig Weichsel. Som det allerede er velkendt fra området, viser foraminiferfaunaerne, at temperaturen generelt var højere i Eem end på noget tidspunkt i Holocæn i det samme område. Derudover viser faunaerne og målinger af stabile isotoper, at det generelt varme klima har været afbrudt af to kuldeperioder. Disse klimatiske svingninger svarer til de, der er beskrevet fra GRIP iskernen på Grønland, og som fornylig også er registreret i terrestriske aflejringer i NV Europa og i dybhavssedimenter i Nordatlanten. Svingningerne skyldes formodentlig variationer i styrke og eventuelt i retning af Den nordatlantiske Strøm (Golfstrømmen).

Aflejringerne fra 178,3–132 m ved Skagen består af ler, sand og grus med et foraminiferselskab bestående af både arktiske og boreale arter. Dette sediment formodes at være afsat under Weichsel Istiden som en bassinudfyldning i denne sydlige udløber af Norske Rende. Lignende sedimenter med blandingsfaunaer, der er præget af store mængder af omlejrede foraminiferer, er også fundet i boringer ved Skærumhede, på Læsø og på Anholt.

De øverste 132 m repræsenterer en marin aflejringsperiode, som dækker Sen Weichsel og hele Holocæn (indtil området blev land for ca. 300 år siden). Foraminiferer og stabile isotoper viser, at der herskede isproximale forhold i området mellem ca. 15.000 og 13.100 BP. I Bølling-Allerød, Yngre Dryas og begyndelsen af Præboreal (13.100–9.600 BP) indikerer den bentoniske fauna relativt koldt, normal salint nordatlantisk vand. Disse atlantiske vandmasser formodes at være strømmet ind langs bunden af Norske Rende som en kompensationsstrøm til udstrømmende smeltevand fra det baltiske område. Svingninger i faunaerne i løbet af perioden tolkes som resultat af, at en reduceret udstrømning af smeltevand i kuldeperioder delvist har opløst stratifikationen af vandsøjlen (Ældre Dryas og Yngre Dryas). Mens de dybere dele af Skagerrak-Kattegat var præget af koldt bundvand helt frem til tidlig Præboreal tid, var lavtvandsområderne periodevist karakteriseret af mildere faunaer allerede fra Bølling Interstadial tid. Først fra omkring 9.600 BP var hele området præget af varme atlantiske

vandmasser, og der herskede fuldt interglaciale forhold overalt.

Ændringer i foraminiferfaunaerne ved Skagen viser, at det nuværende strømsystem i Nordsøen og Skagerrak-Kattegat området må være blevet etableret ca. 7.600 BP. Dette kobles sammen med åbningen af Den engelske Kanal og de danske bæltter. Den videre faunaudvikling og ændringen i lithologi tyder desuden på en mindre modifikation af cirkulationssystemet omkring 5.500 BP, hvilket sandsynligvis er et resultat af den generelle afkøling i Nordatlanten ved slutningen af det Holocæne klimatiske optimum.

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